

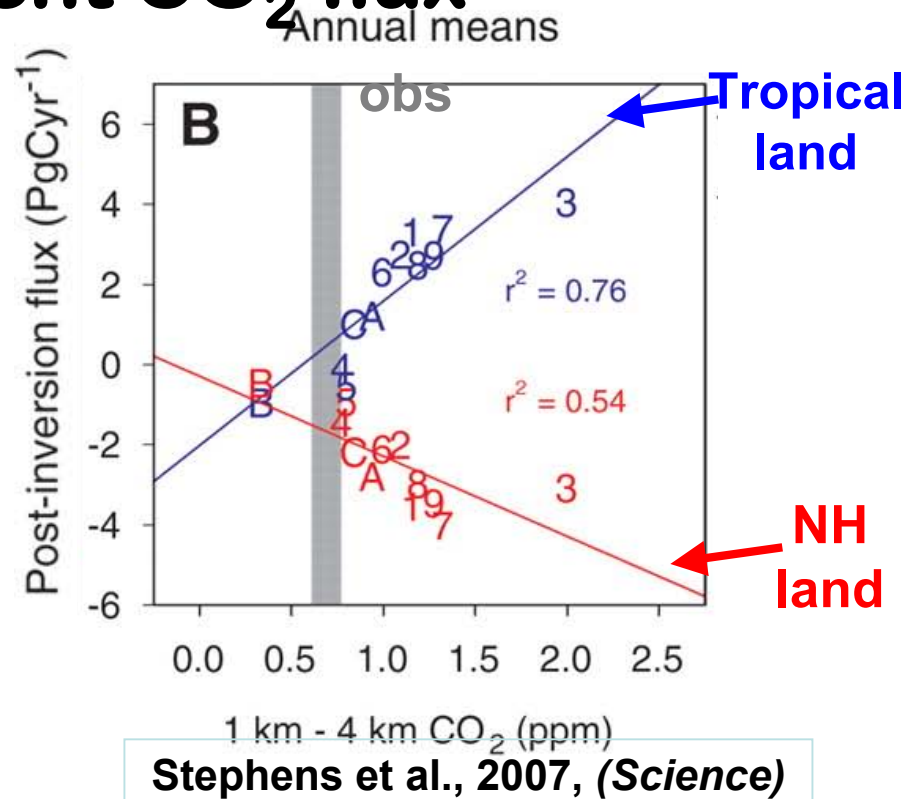
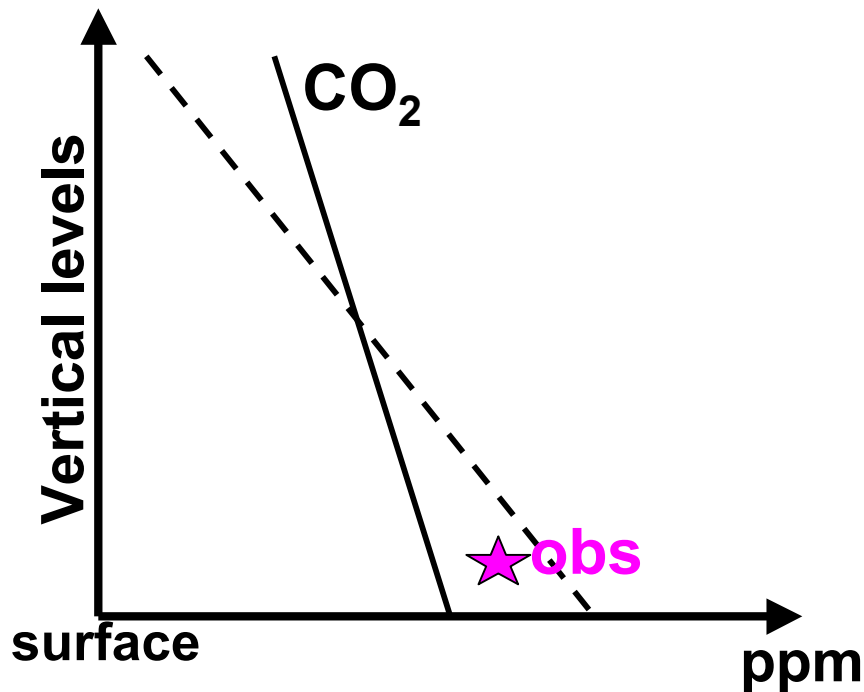
Assimilation of AIRS CO₂ Observations with EnKF in a Carbon-Climate Model

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Mous Chahine, Ed Olsen, Luke Chen (JPL)

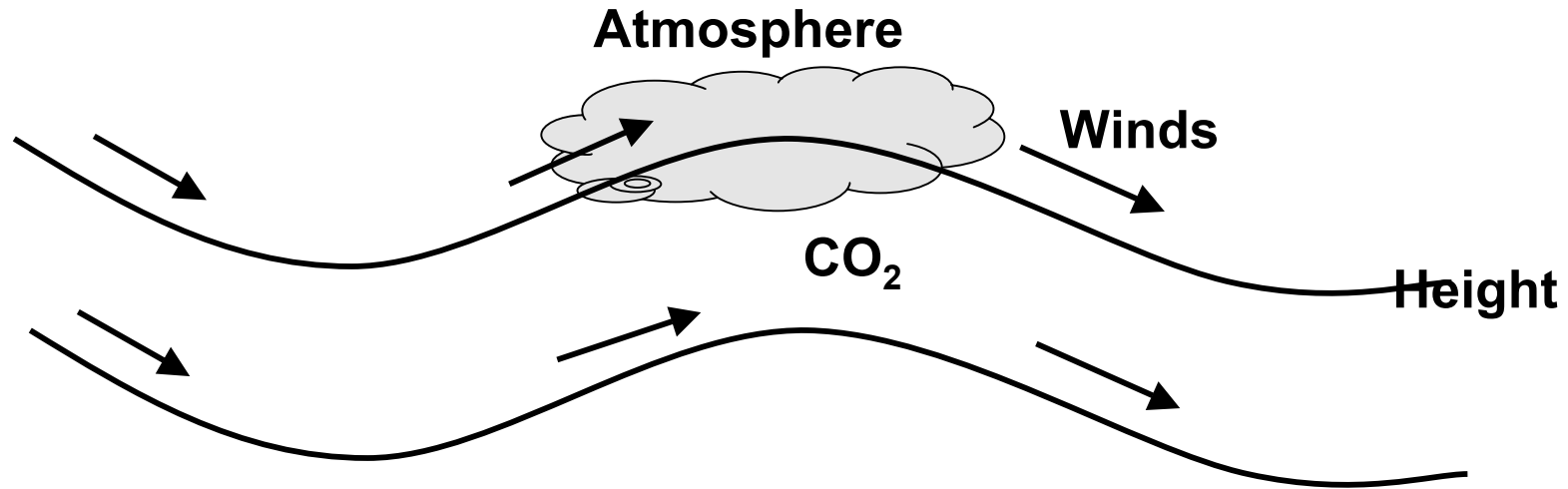
Different CO₂ vertical gradient forecasts give different CO₂ flux



- Stephens et al., 2007, (*Science*)
- Numbers and characters are different transport models.

=> Important to have accurate vertical mixing in the model;
=> Accurate 4-D (x, y, z, t) CO₂ fields.

Meteorology fields & CO₂



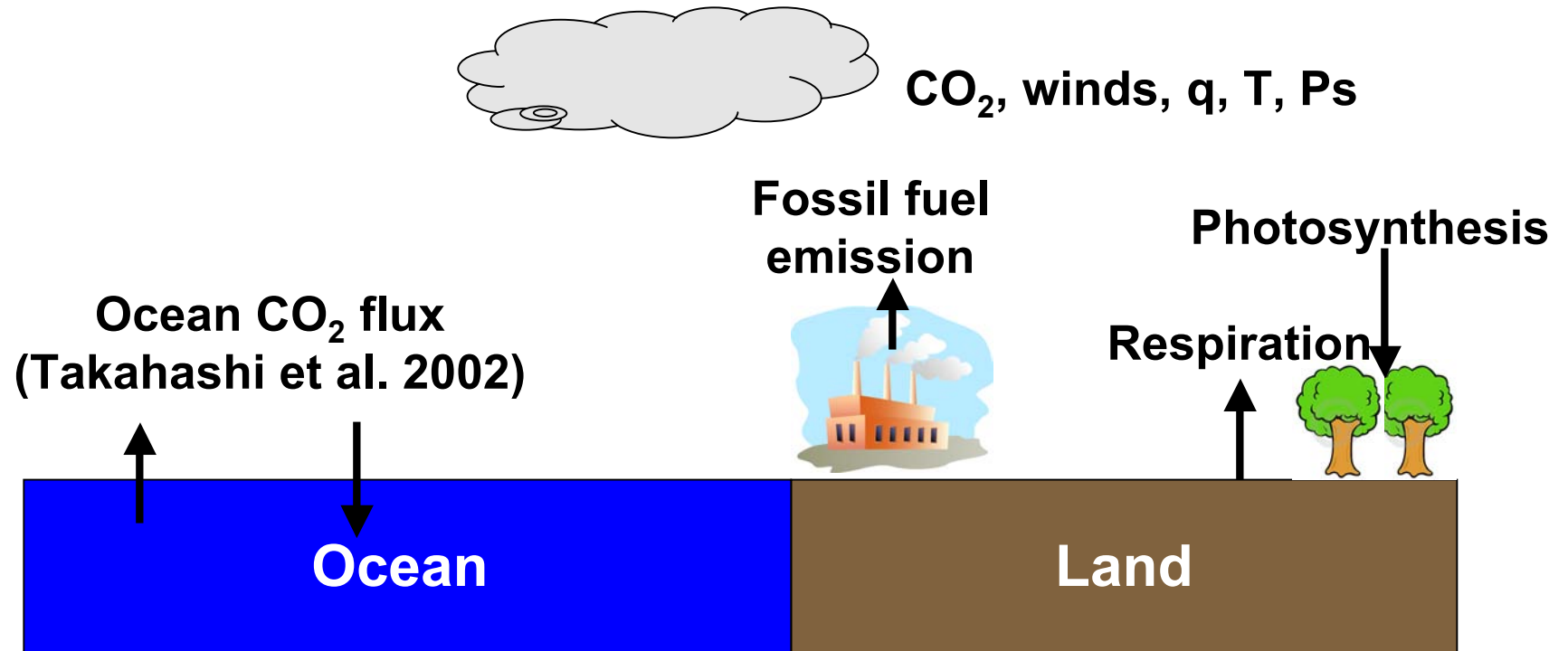
- Offline transport models have been used.
- The initialization meteorology fields are from either reanalysis products (usually 6-hourly) or off-line dynamical model;
- The vertical mixing has large uncertainty;
- Single realization of meteorological field.

Research Goals

- **Generate 6-hourly 3-D (x, y, z) CO₂ fields by assimilating CO₂ and meteorological observations with full GCM**

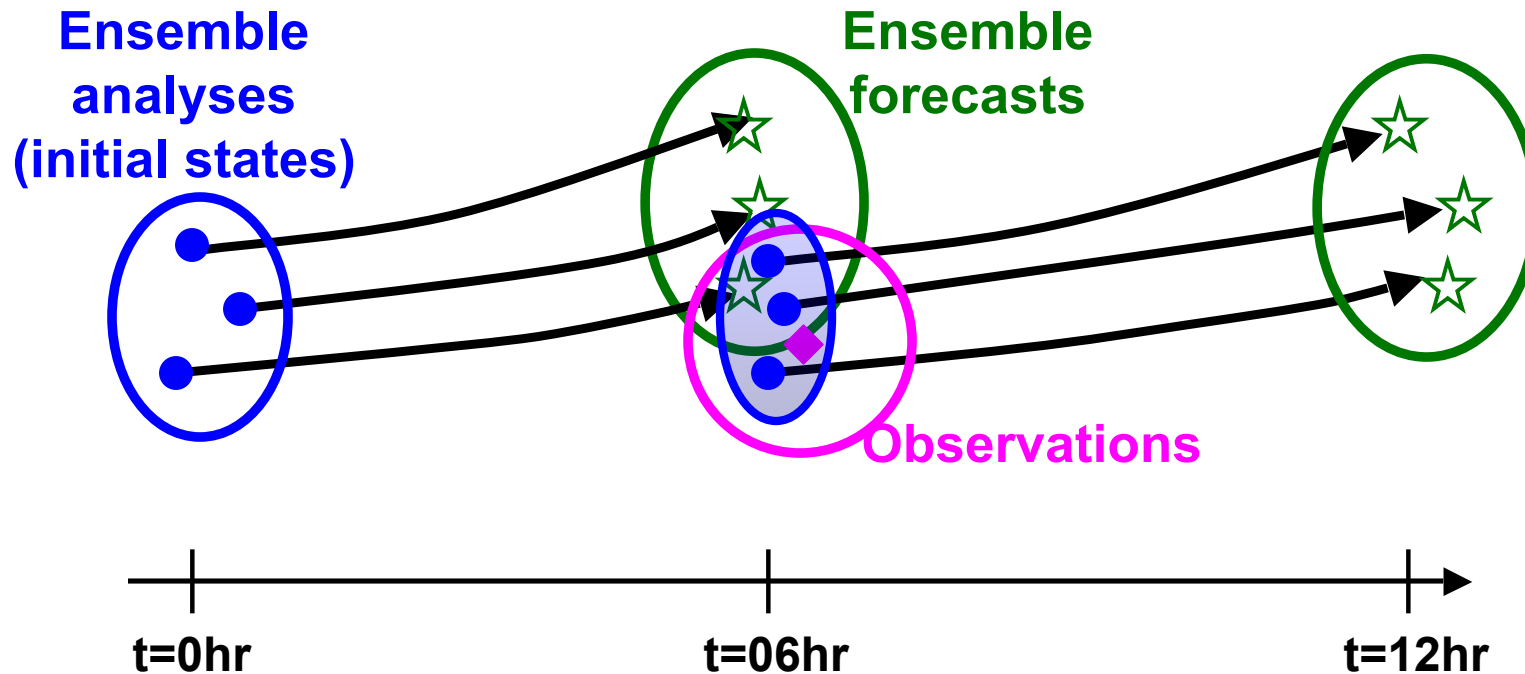
Carbon-Climate Model

Community Atmospheric Model
(fvCAM 3.5) (2.5x1.9x26)



- CO_2 is transported as a tracer;
- Vertical mixing is updated every 30 minutes;
- Land carbon flux: 6-hourly flux from biogeochemical model.₅

Ensemble Kalman Filter (EnKF) process



- Forecast error changes with time;
- Obtain ensemble analyses.

CO₂ Observation Operator

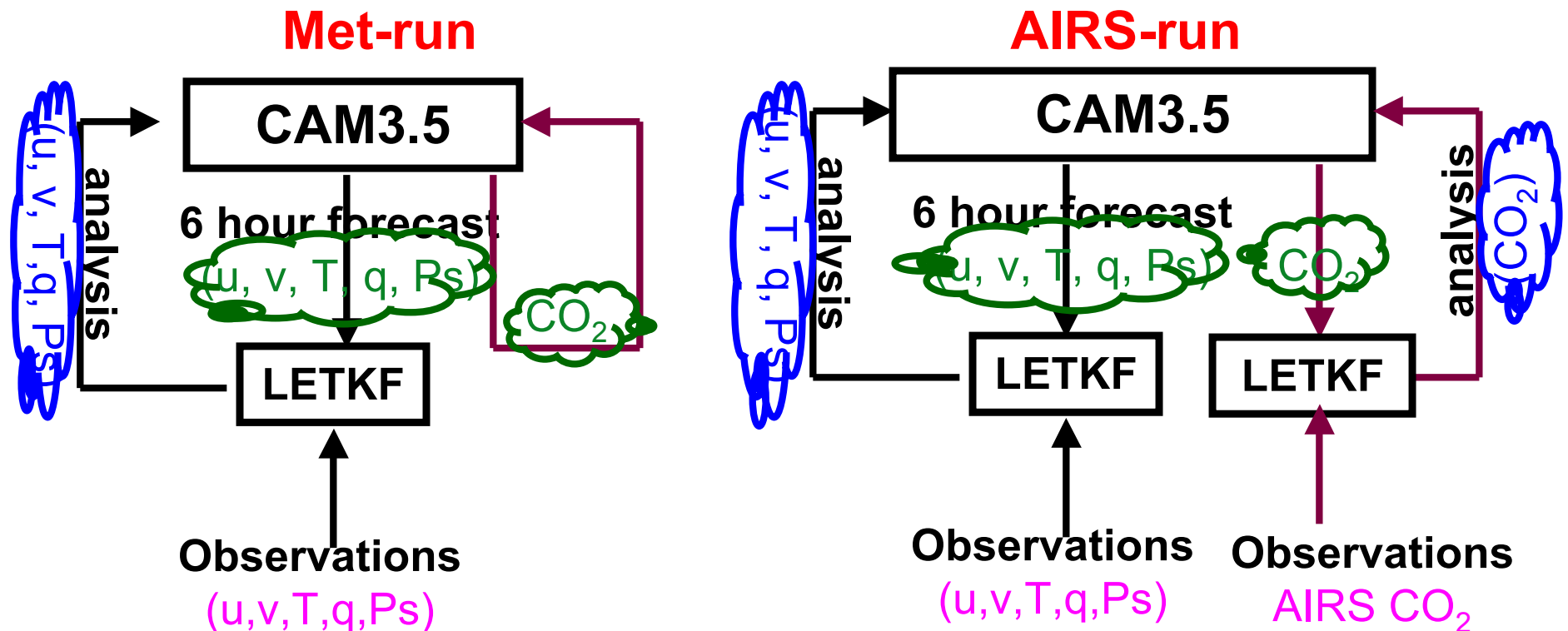
- Model forecast x^b is CO₂ vertical profile;
 - AIRS CO₂ is column-weighted Volume Mixing Ratio (vmr);
- => **observation operator**: interpolate x^b to obs location & calculate model forecast column-weighted CO₂ vmr.

$$\underbrace{\mathbf{y}^b}_{\text{model forecast "obs"}} = \underbrace{\underbrace{\mathbf{A}^T}_{\text{avg kernel}} \left(\underbrace{\mathbf{S}}_{\text{spatial interpolator}} \right)}_{\text{obs operator}} \left(\underbrace{\mathbf{x}^b}_{\text{model forecast}} \right)$$

Assimilation experiments

- **Met-run: assimilate raw meteorological observations (10^6 observations)**
- **AIRS-run: assimilate AIRS CO₂ observations in conjunction with meteorological observations.**

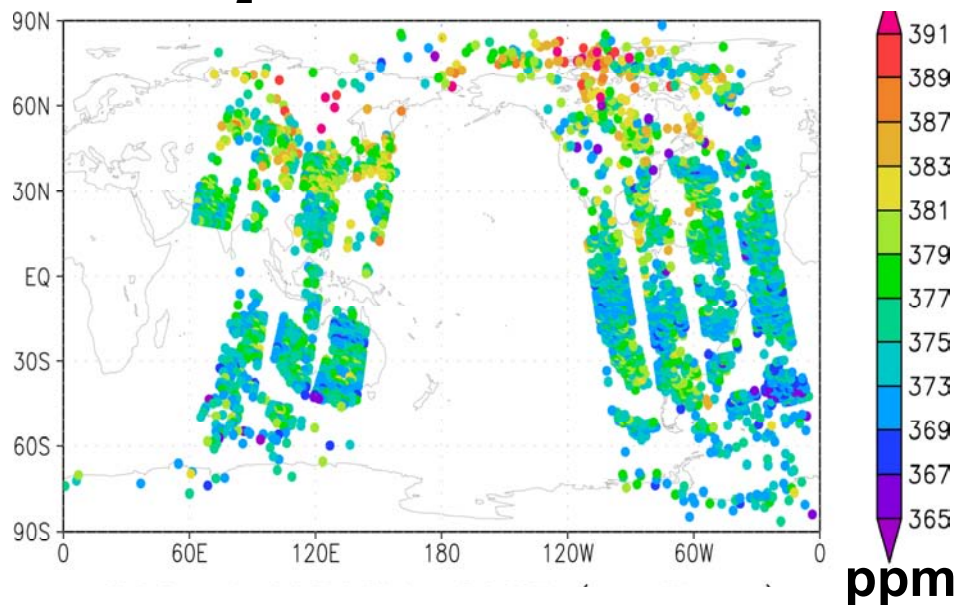
The impact of AIRS CO₂ assimilation on 6-hourly CO₂ 3D (x, y, z) fields



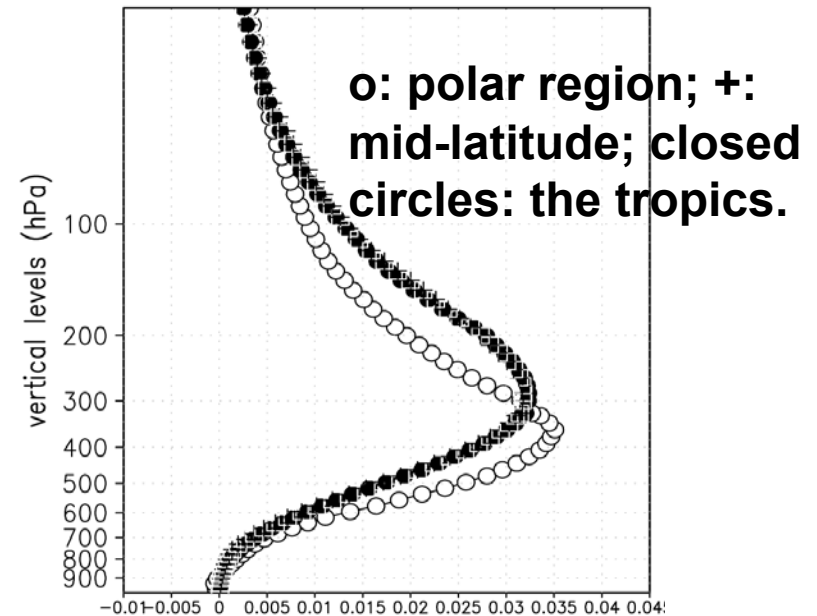
- AIRS-run: AIRS CO₂+met obs; Met-run: only met obs.
- The year of 2003.
- Prescribed surface CO₂ flux forcing.

More than 2000 AIRS CO₂ within 6 hours; more sensitive in the middle troposphere

AIRS CO₂ at 18Z01May2003 (+/-3hour)



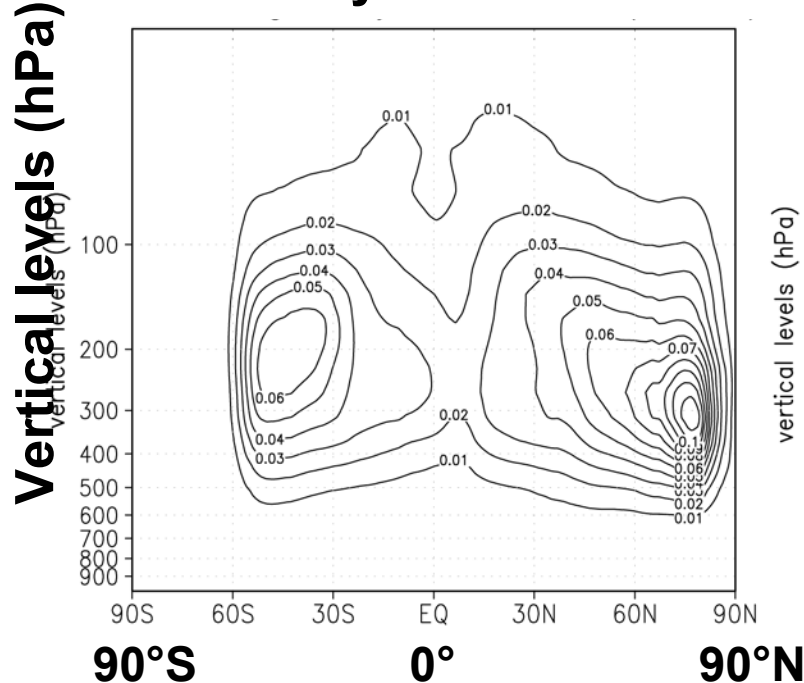
AIRS averaging kernel



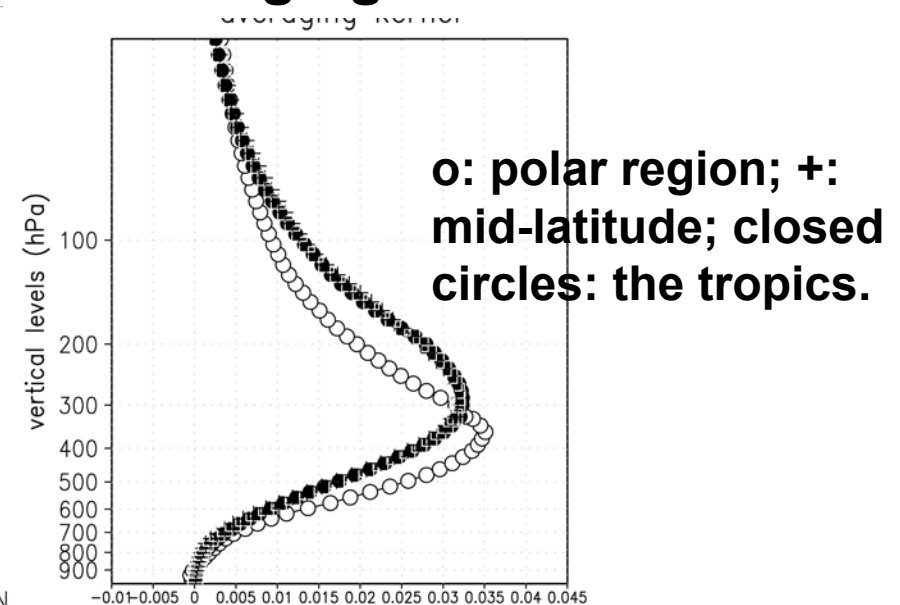
- Averaging kernel: the sensitivity of AIRS CO₂ to CO₂ at each vertical level.

Analysis corrections to CO₂ forecast peak at the similar levels as the peak of the averaging kernels

Time-averaged (10 months)
absolute analysis corrections



Averaging Kernel

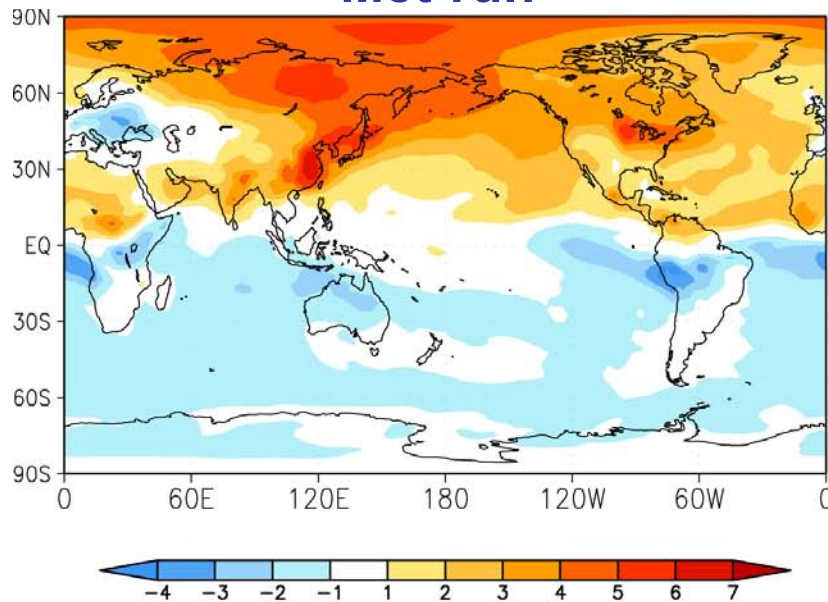


- No CO₂ observations beyond 60°S.

Assimilating CO₂ adjusts CO₂ vertical gradient

May 2003: CO₂(850hPa)-CO₂(400hPa)

Met-run

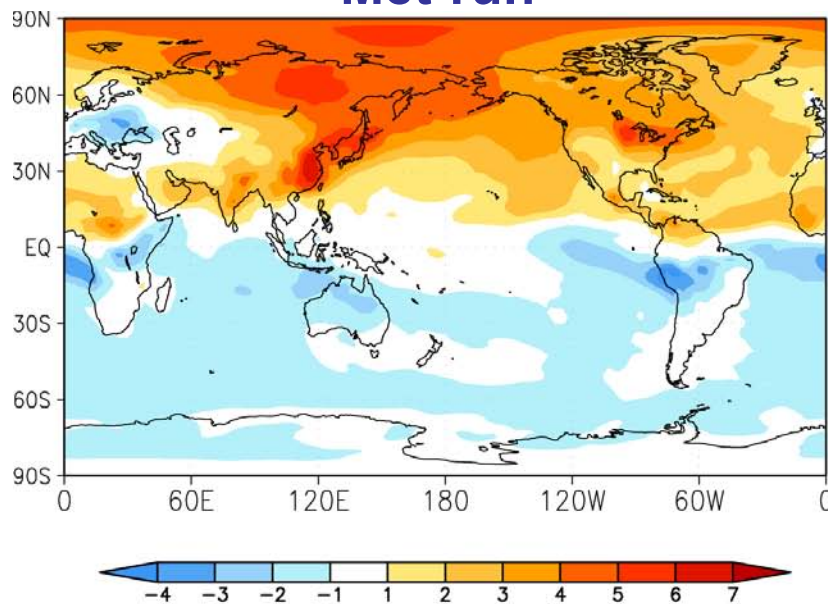


- In the NH, CO₂(850hPa) > CO₂(400hPa): fossil fuel+ land carbon source;
- In the SH, CO₂(850hPa) < CO₂(400hPa): transported from the NH.

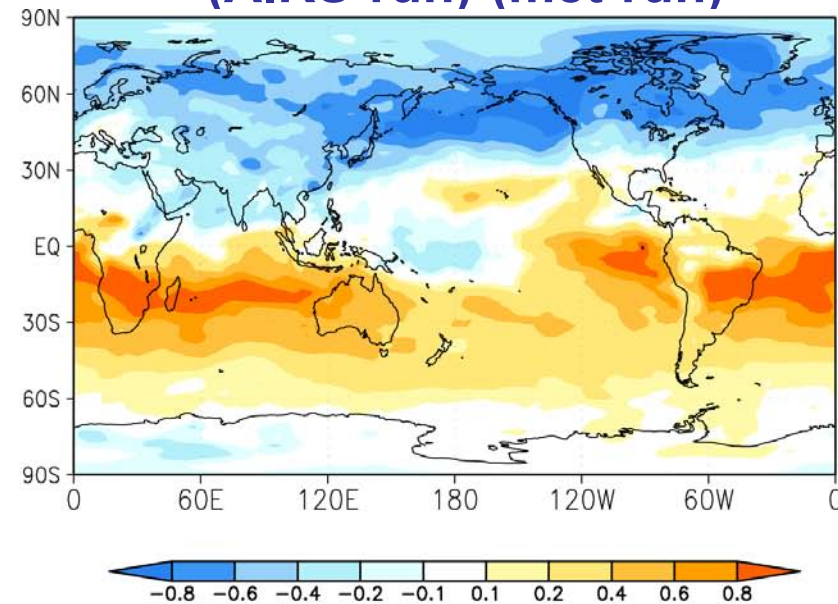
Assimilating CO₂ adjusts CO₂ vertical gradient

May 2003: CO₂(850hPa)-CO₂(400hPa)

Met-run



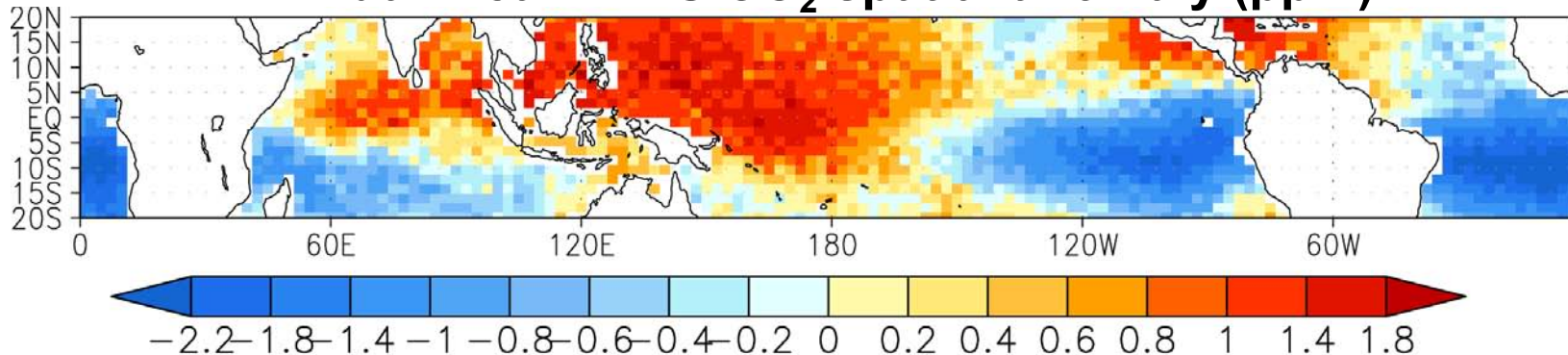
(AIRS-run)-(met-run)



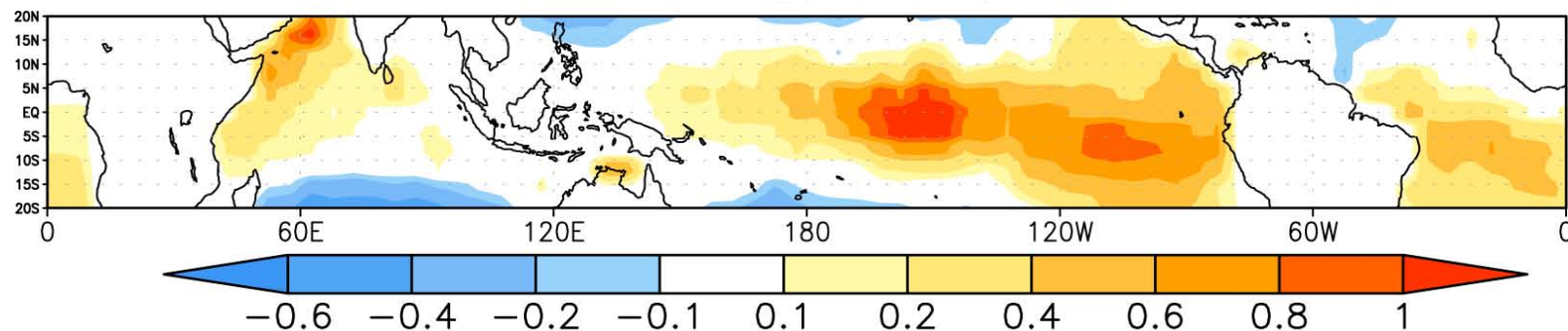
- In the NH, CO₂(850hPa)>CO₂(400hPa): fossil fuel+ land carbon source;
- In the SH, CO₂(850hPa)<CO₂(400hPa): transported from the NH.
- Require CO₂ obs in the lower troposphere to further constrain gradient.

Inconsistent spatial distribution between AIRS CO₂ and ocean-air CO₂ flux

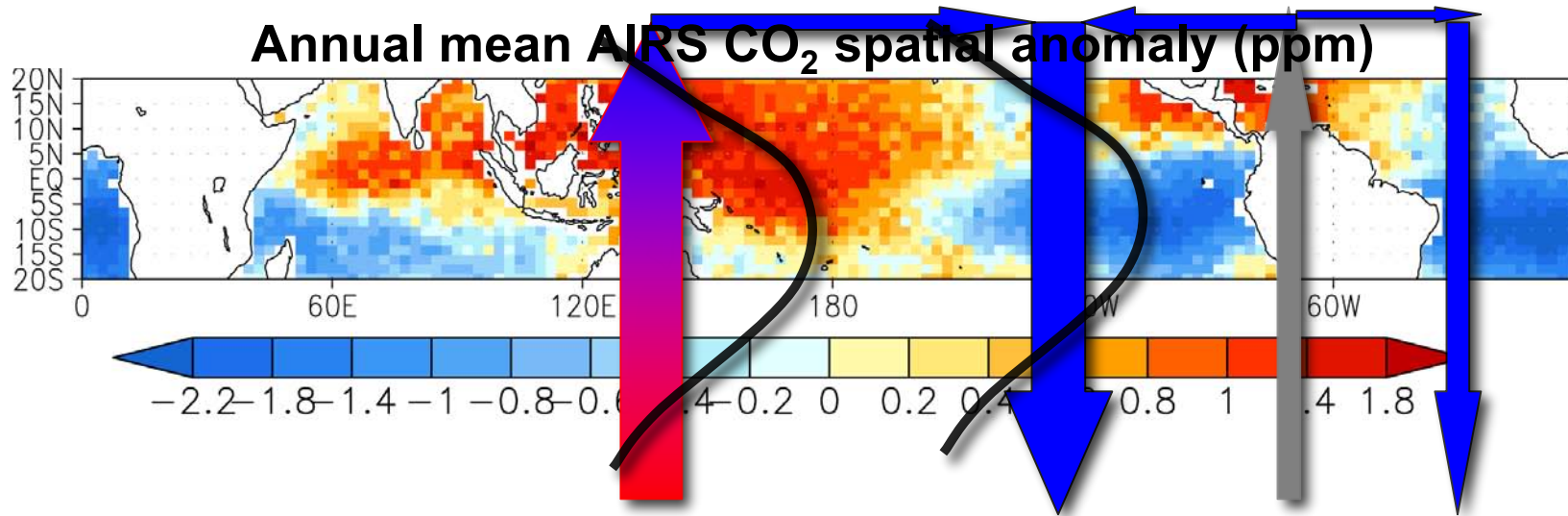
Annual mean AIRS CO₂ spatial anomaly (ppm)



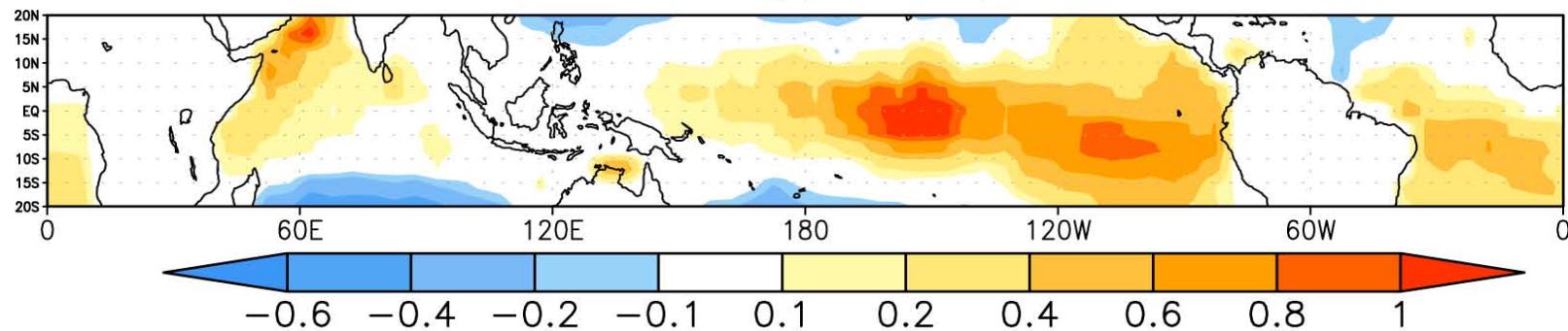
Ocean-atmosphere CO₂ flux (unit: 10⁻⁹kgC/m²/s Takahashi et al., 2002)



Tropical AIRS CO₂ relates to circulation and averaging kernel

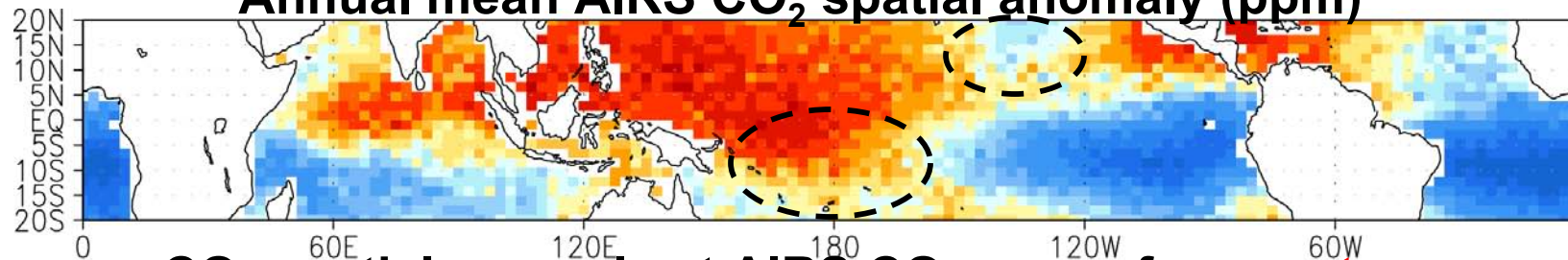


Ocean-atmosphere CO₂ flux (unit: 10⁻⁹kgC/m²/s Takahashi et al., 2002)

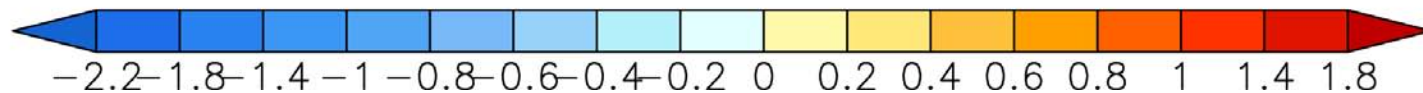
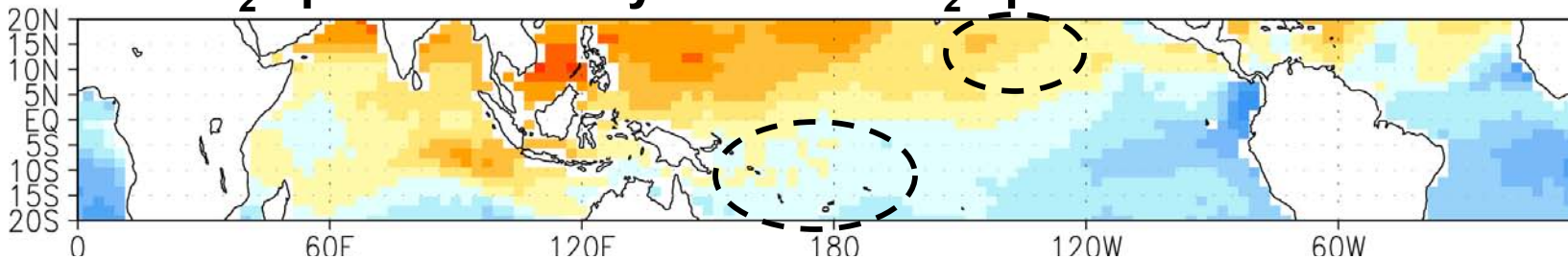


Assimilating AIRS CO₂ improves spatial pattern

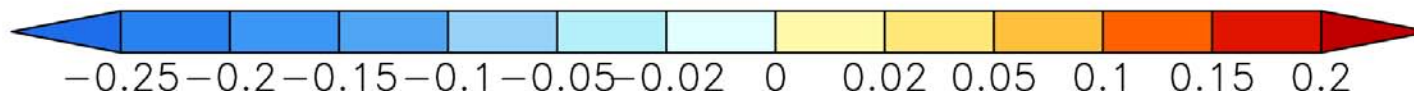
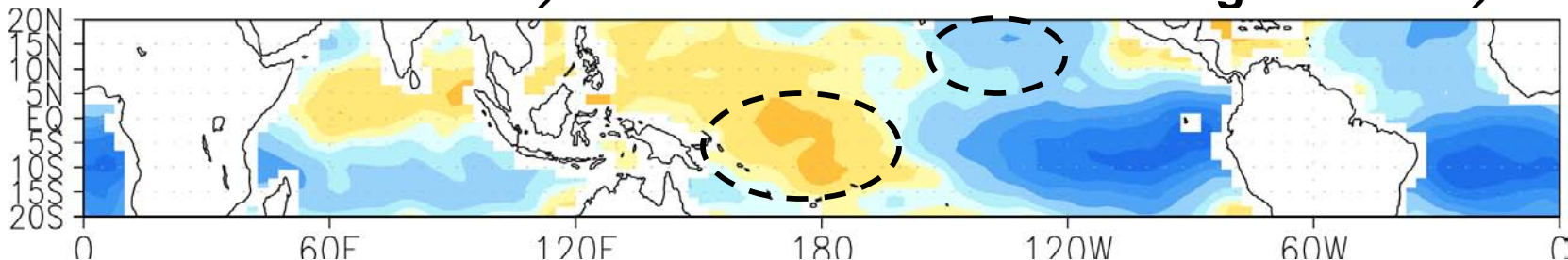
Annual mean AIRS CO₂ spatial anomaly (ppm)



CO₂ spatial anomaly at AIRS CO₂ space from **met-run**



Annual mean CO₂ correction from assimilating AIRS CO₂

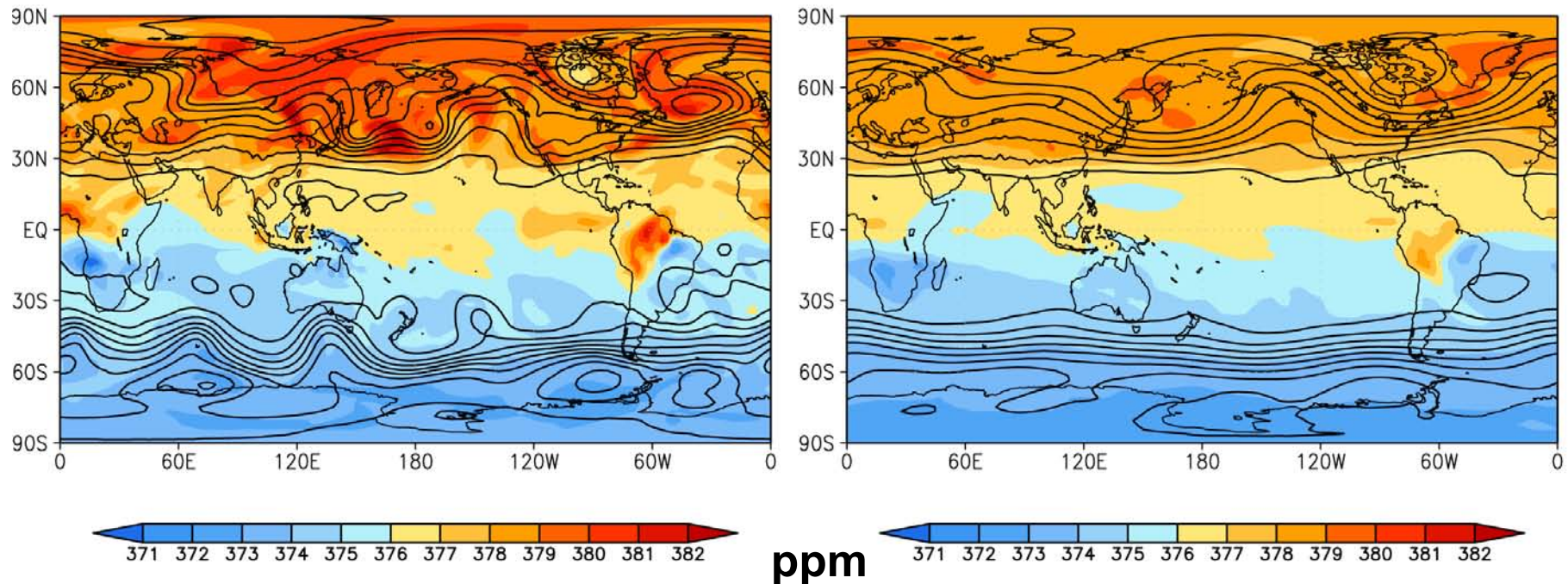


Consistent CO₂ distribution and weather pattern

500hPa geopotential height (contour) and CO₂ from AIRS-run (shaded)

Single time (12Z27Feb2003)

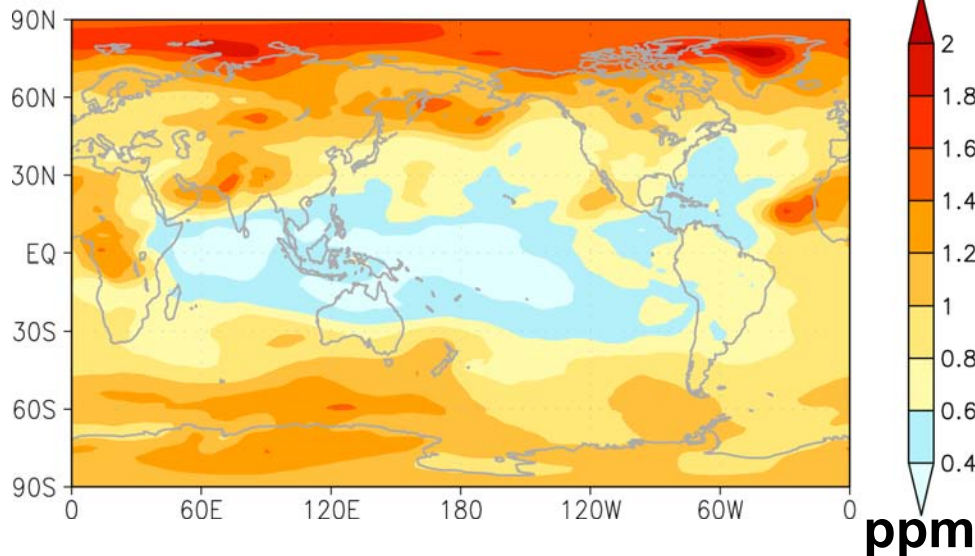
Time average over Feb 2003



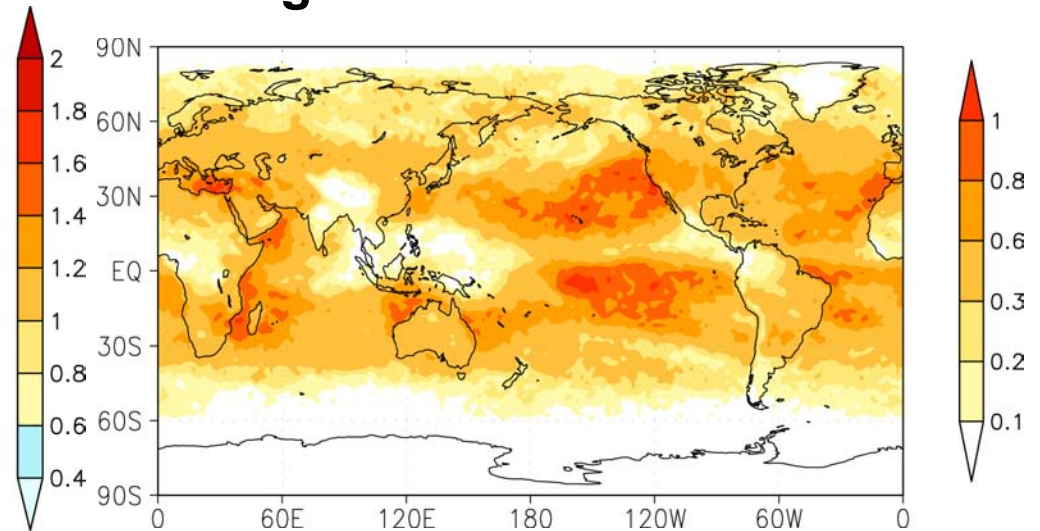
- Simultaneous assimilation of meteorology variables and CO₂ generates CO₂ distribution consistent with weather pattern

CO₂ analysis spread ranges from 0.4ppm to 2ppm at 400hPa

400hPa monthly mean
(September) CO₂ spread



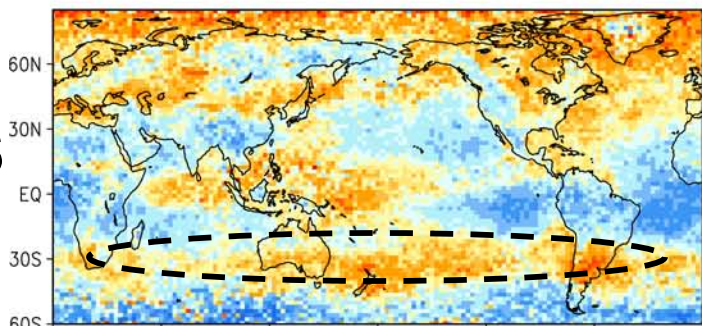
Average num of CO₂ observations at
each grid box within 6 hours



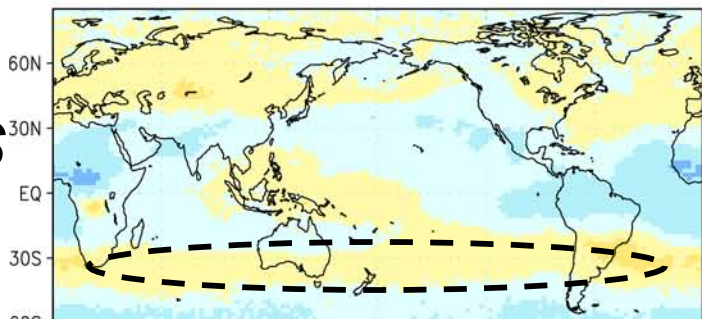
- Analysis ensemble spread is related to observation coverage, forecast error and observation error;
- Larger spread over high latitudes, and over land;
- Smaller spread over tropical ocean is due to observation coverage and propagation through forecast.

Column-integrated CO₂—Sep

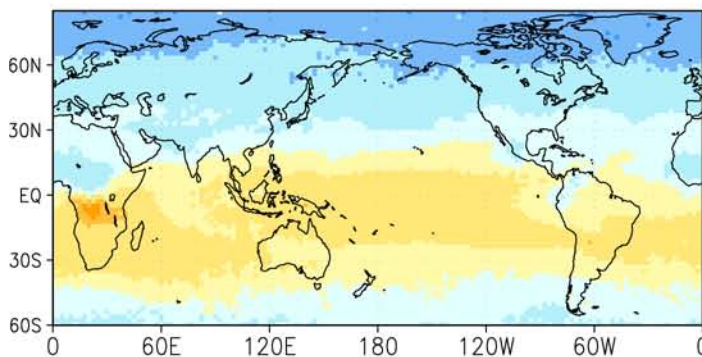
AIRS



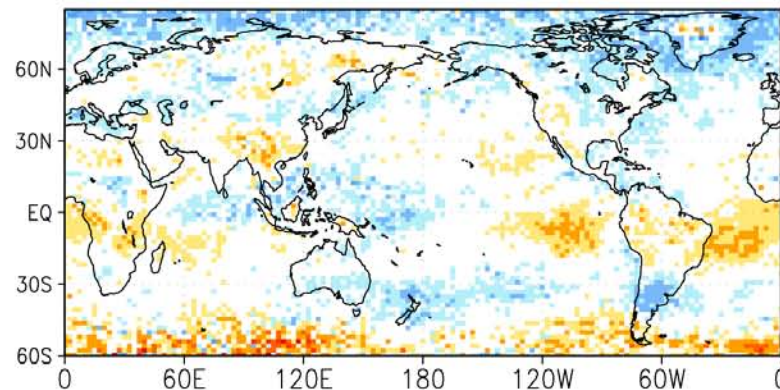
**AIRS
-run**



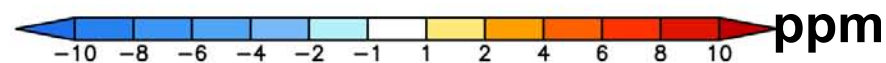
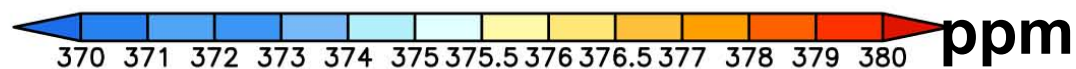
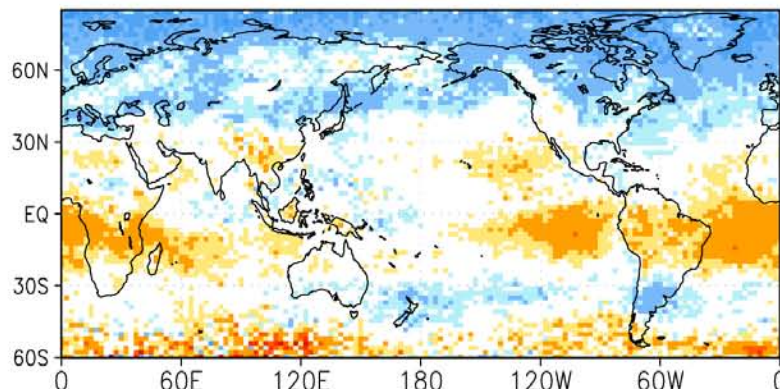
**Met-
run**



AIRS-run - AIRS

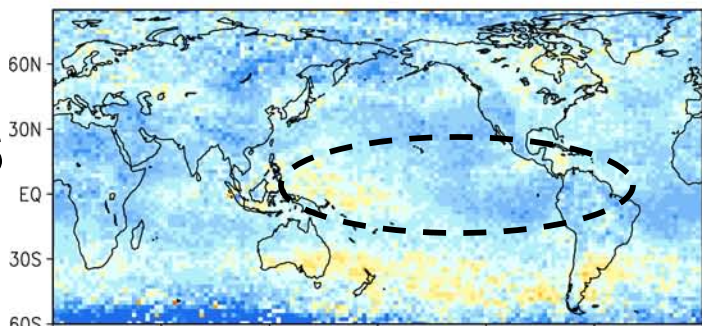


Met-run - AIRS

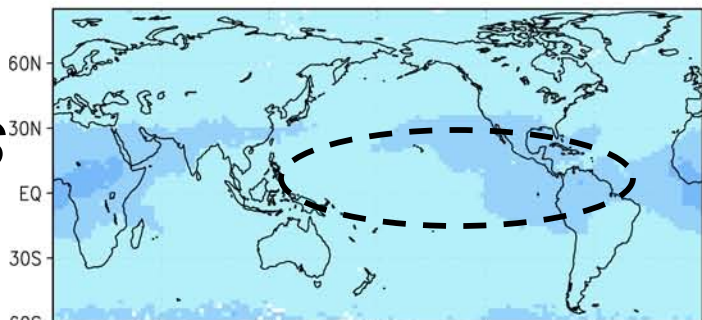


Column-integrated CO₂—Oct

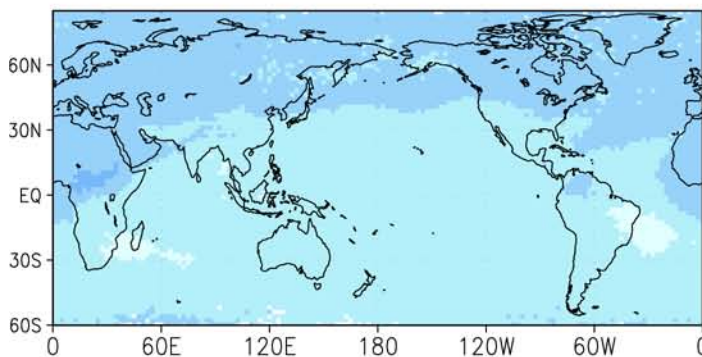
AIRS



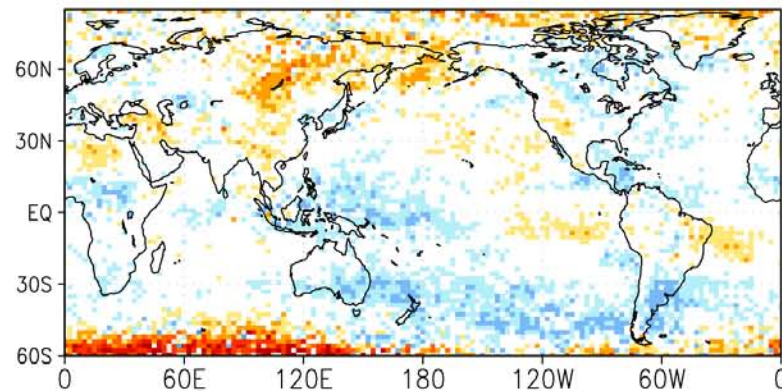
**AIRS
-run**



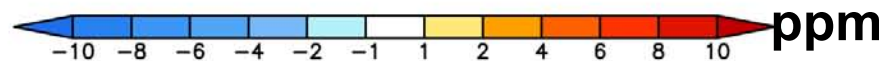
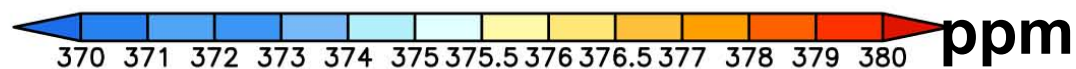
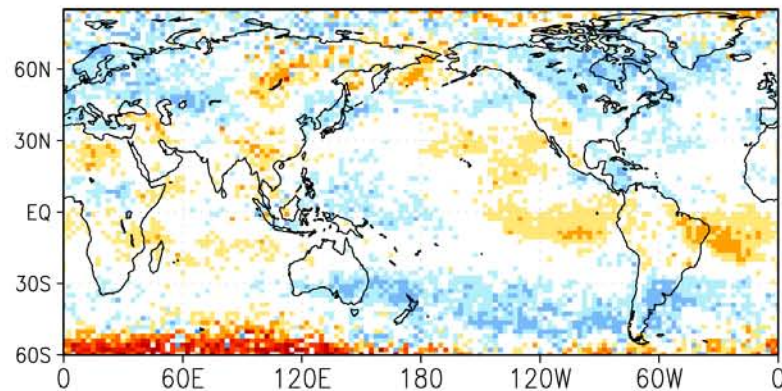
**Met-
run**



AIRS-run - AIRS

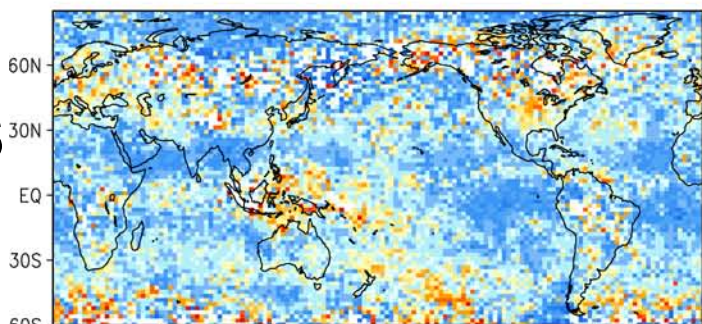


Met-run - AIRS

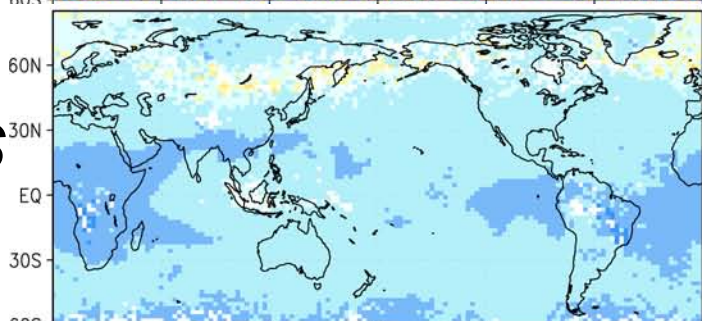


Column-integrated CO₂—Nov

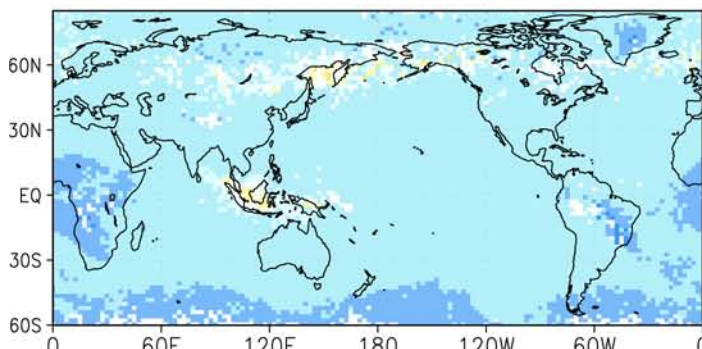
AIRS



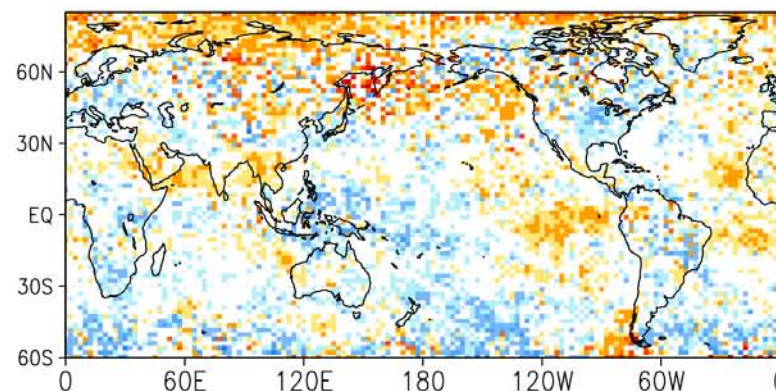
**AIRS
-run**



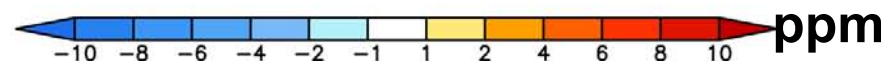
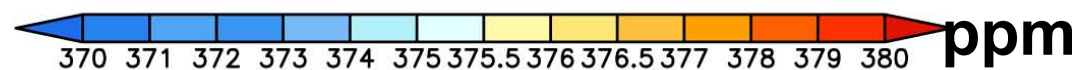
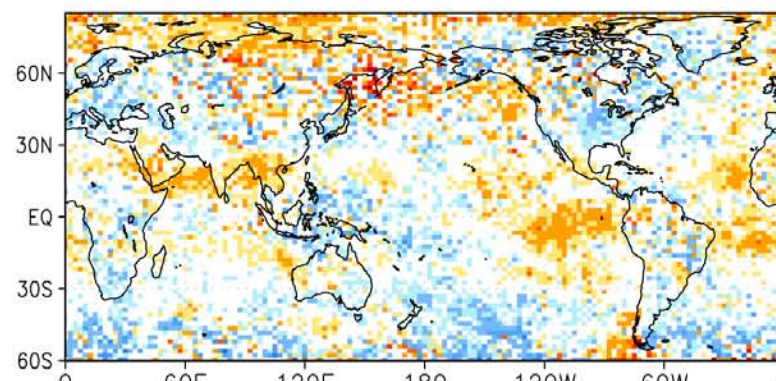
**Met-
run**



AIRS-run - AIRS



Met-run - AIRS

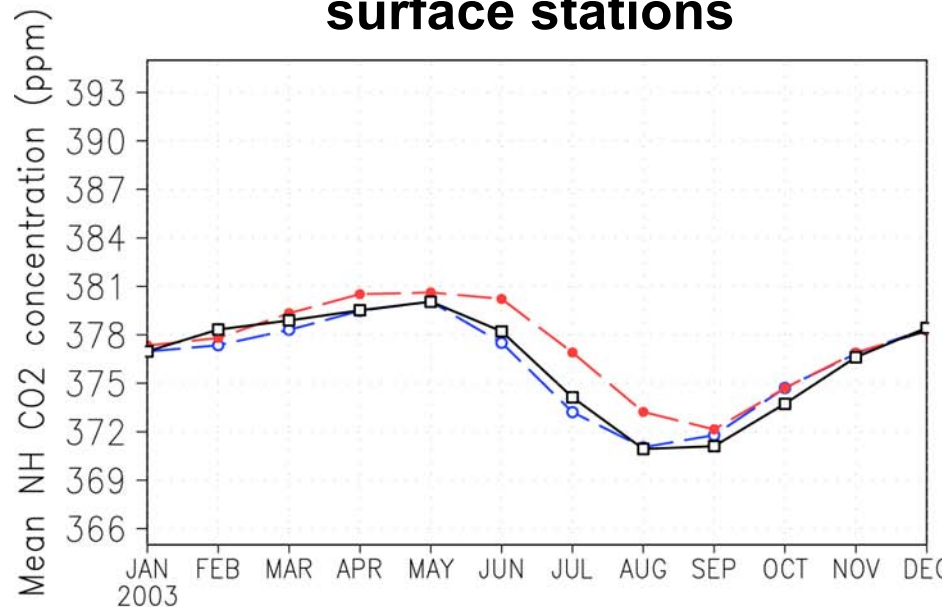


**20 days with no AIRS
CO₂**

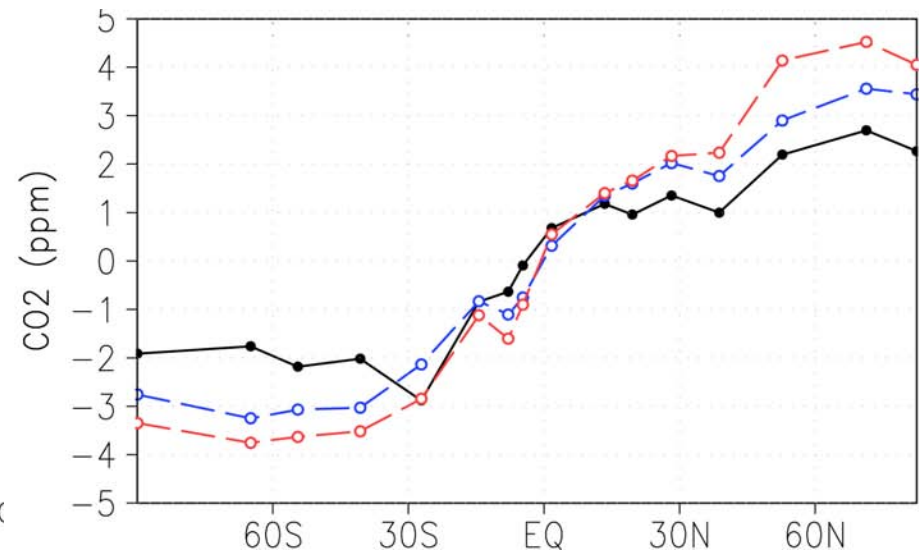
Assimilating AIRS CO₂ improves surface CO₂ seasonal cycle and the N-S gradient

Surface obs: black; Met-run: red; AIRS-run: blue

Mean NH CO₂ concentration at 8 surface stations



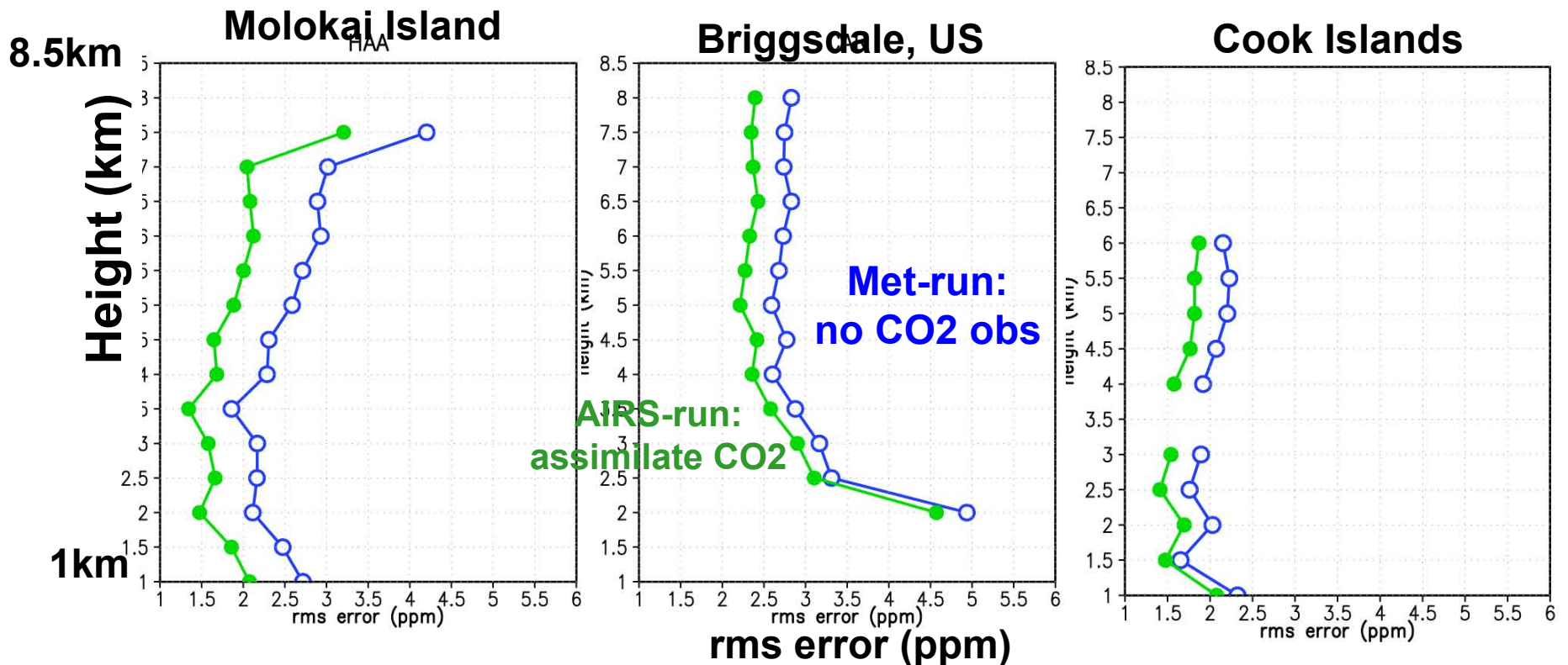
The N-S gradient based on 16 surface stations



Surface data is from NOAA/ESRL website

Met-run has similar NH CO₂ concentration and the N-S gradient

Assimilating AIRS CO₂ improves CO₂ state estimate

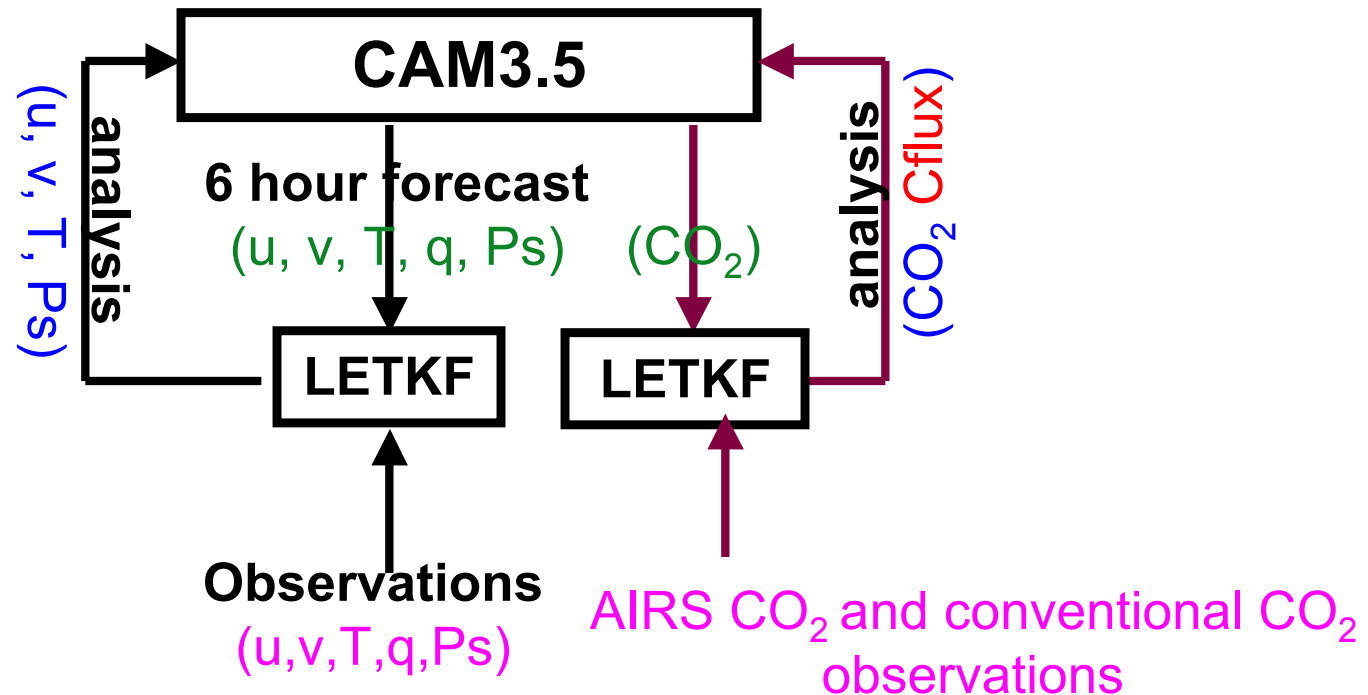


Verified against independent aircraft CO₂ observations

- CO₂ from the AIRS-run can be about 1 ppm more accurate than those from the met-run.

Preliminary results on surface carbon flux estimation by assimilating AIRS CO₂

The impact of AIRS CO₂ assimilation on surface CO₂ flux

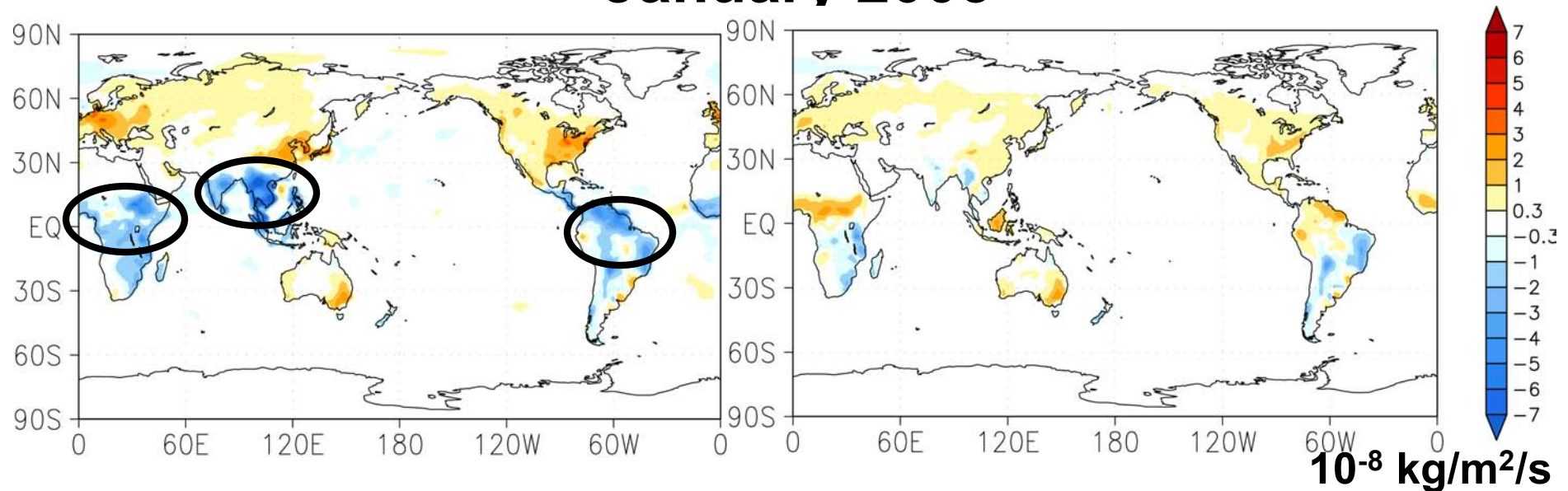


LETKF: Local Ensemble Transform Kalman Filter (Hunt et al., 2007)

- The carbon flux analysis acts as boundary forcing for the forecast of next time step.
- Four and a half months assimilation cycles (01Jan2003-10May2003).

Carbon Flux Analysis: Data Assim (left) Carbon Flux (CASA (land)+Takahashi (ocean))(right)

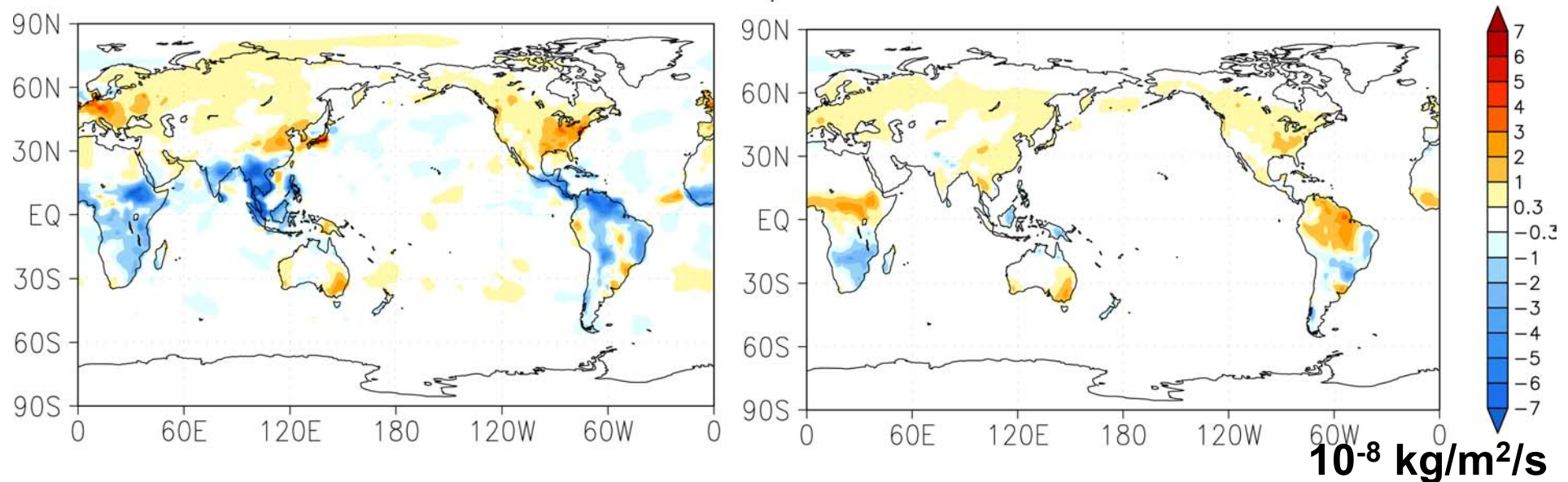
January 2003



- AIRS has the most impact over the tropical land
- Stronger source in the NH winter
- Stronger sink in the tropics and SH subtropics
- Noisy over ocean compared to Takahashi

Carbon Flux Analysis: Data Assim (left) Carbon Flux (CASA (land)+Takahashi (ocean))(right)

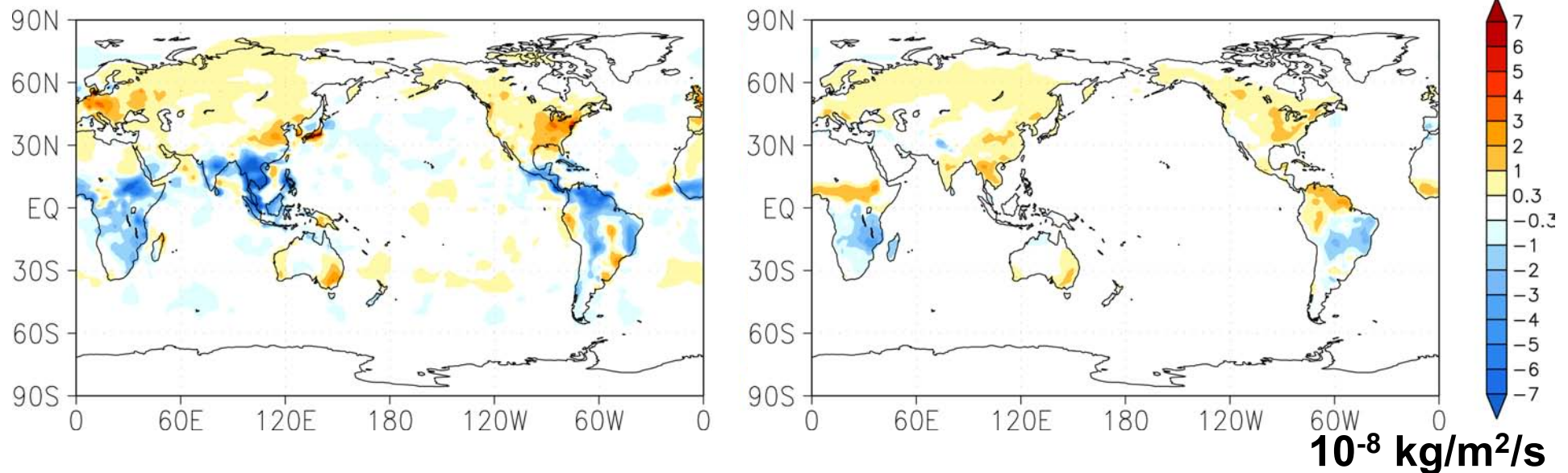
February 2003



- **Stronger source** in the NH winter
- **Stronger sink** in the tropics and SH subtropics
- **Noisy** over ocean compared to Takahashi

Carbon Flux Analysis: Data Assim (left) Carbon Flux (CASA (land)+Takahashi (ocean))(right)

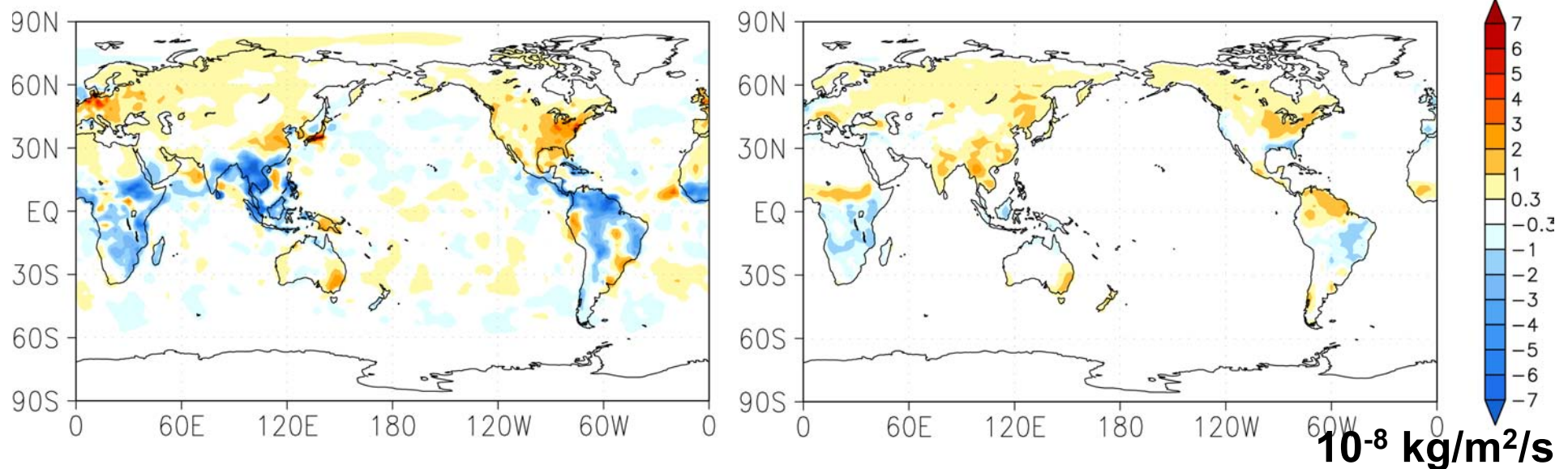
March 2003



- Little change

Carbon Flux Analysis: Data Assim (left) Carbon Flux (CASA (land)+Takahashi (ocean))(right)

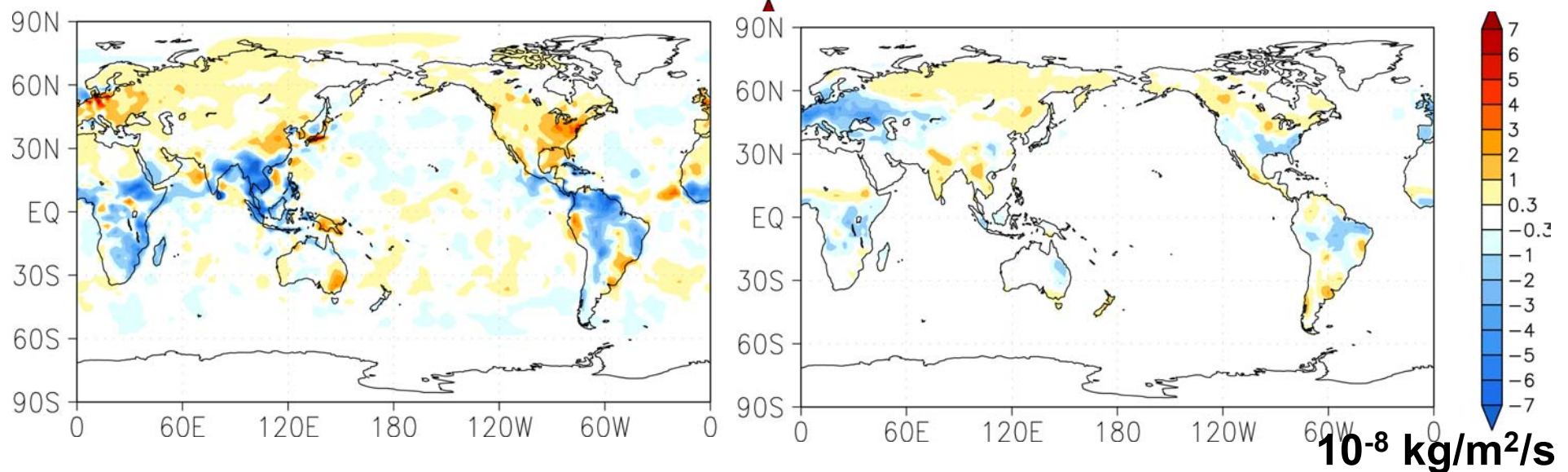
April 2003



- Little change

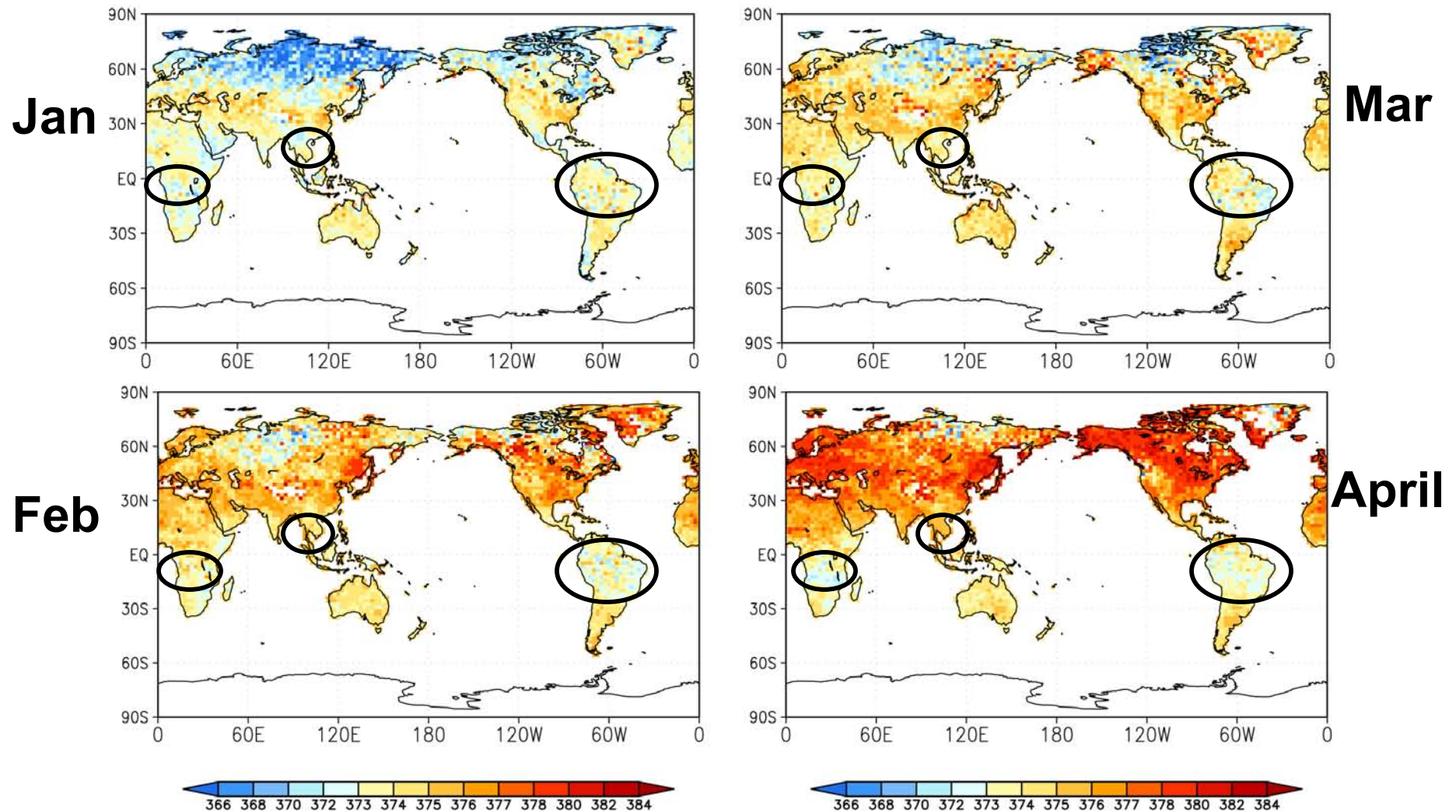
Carbon Flux Analysis: Data Assim (left) Carbon Flux (CASA (land)+Takahashi (ocean))(right)

May 2003



- As in the OSSEs, the surface fluxes appear initially to be reasonable and then they “get stuck”.

Monthly average AIRS CO₂ does not change much over the tropical land from January to May



Conclusions - General

- **EnKF brings important advantages for Reanalysis:**
 - ✓ **Analysis uncertainty**
 - ✓ **Adaptation to new observing systems**
 - ✓ **Estimation of obs. errors and identification of bad observations (not shown)**
 - ✓ **Estimation of model bias (essential)**
- **For Carbon Reanalysis it is essential to assimilate at the same time meteorological and carbon observations**

Conclusions - Simulations (OSSEs)

- The advantage of the OSSEs is that we know the true fluxes and CO₂
- It is possible to estimate surface carbon fluxes from atmospheric CO₂ measurements but
 - ✓ Need “variable localization” to reduce sampling errors
 - ✓ Need adaptive inflation of the B error covariance
 - ✓ Need to estimate model bias
- Problem: the initial results after spinup from random fields are good, but then the surface fluxes “don’t want to change anymore”.
 - ✓ This is probably due to model bias

Conclusions - AIRS data assimilation

- **AIRS CO2 data assimilation is clearly successful!**
 - ✓ **Improved atmospheric CO2 and N-S gradient**
 - ✓ **Better agreement with independent observations**
 - ✓ **Insight about vertical circulation and mixing**
- **Preliminary estimations of carbon fluxes are very promising after one month spin-up:**
 - ✓ **Compared with CASA fluxes they yield reasonable uptakes in the SH summer and stronger emissions in the NH winter**
 - ✓ **But, like in the OSSEs, the fluxes “don’t want to change” with season**

Conclusions - AIRS/IASI GoSAT/OCO2

- The combination of satellite and in situ data is important:
 - ✓ Results are more accurate in NH than in SH
- We need more near surface information
 - ✓ Carbon fluxes can be derived from atmospheric CO2
- Results depend on optimal forecast spread, a difficult problem for surface fluxes:
 - ✓ Work on estimating model bias
 - ✓ Should find why after initial good surface fluxes they “don’t want to change” with season
 - ✓ We probably need to estimate diurnal and seasonal changes with a different approach (e.g., EOFs).³⁵